

blast and the M 2.8 Paso Robles, CA earthquake. The optimized model yielded a crustal root beneath the Sierra Nevada Mountains with an approximate depth of 55 km, assuming a Moho velocity of 7.6 km/s beneath the Sierra Nevada, and centered at least 100 km west of the topographic crest (Pullammanappallil and Louie, 1994). To the northeast of the root, the inversion confirms a typical 30-32 km-thick crust in northern Nevada, thickening to the north toward the Nevada-Idaho state border. The model agrees with the 1986 PASSCAL refraction and the COCORP 40N reflection experiment that both revealed 30 km thick crust in northern Nevada (Klemperer et al., 1986; PASSCAL Working Group, 1988; Benz et al., 1990).

In order to confirm and expand on the experiment's basic model, a second tomography inversion was completed (Figure 2). The second model utilized re-picked first arrivals from the Paso Robles, CA earthquake and the Barrick GoldStrike 232 blast, as well as an additional Barrick GoldStrike blast on Julian day 233. The second inversion confirms our original picks (without a 2.5 second delay that was originally added to the Paso Robles, CA event), and adds clarification to features in the initial version (Pullammanappallil and Louie, 1994). From Model 2, we observe a crustal root of only 44 km +/- 2 km with steep sides, centered 40 km west of the topographic crest at the latitude of approximately 37 N. Within northern Nevada, we observe 30-32 +/- 2 km-thick crust that is in agreement with Model 1, as well as the 1986 PASSCAL and the COCORP 40 N transect (Klemperer et al., 1989; Benz et al., 1990; PASSCAL Working Group, 1988). This gives confidence to the picks and the model inversion. There is little transition between the steep-sided root on its eastern side and the Basin and Range. At the latitude of Long Valley, we observe a steeply dipping 14-20 km step in the Moho. Assuming a root Pn velocity of 7.6 km/s and a Pn velocity in Nevada of 7.8 km/s, this forces the Moho interpretation of Model 2 to cut across the velocity boundaries defined by the inversion. Where the change from 7.6 to 7.8 km/s actually takes place cannot be determined from this study alone. We



Figure 2. Updated model for the INC seismic refraction transect (see Figure 1 for INC transect location).

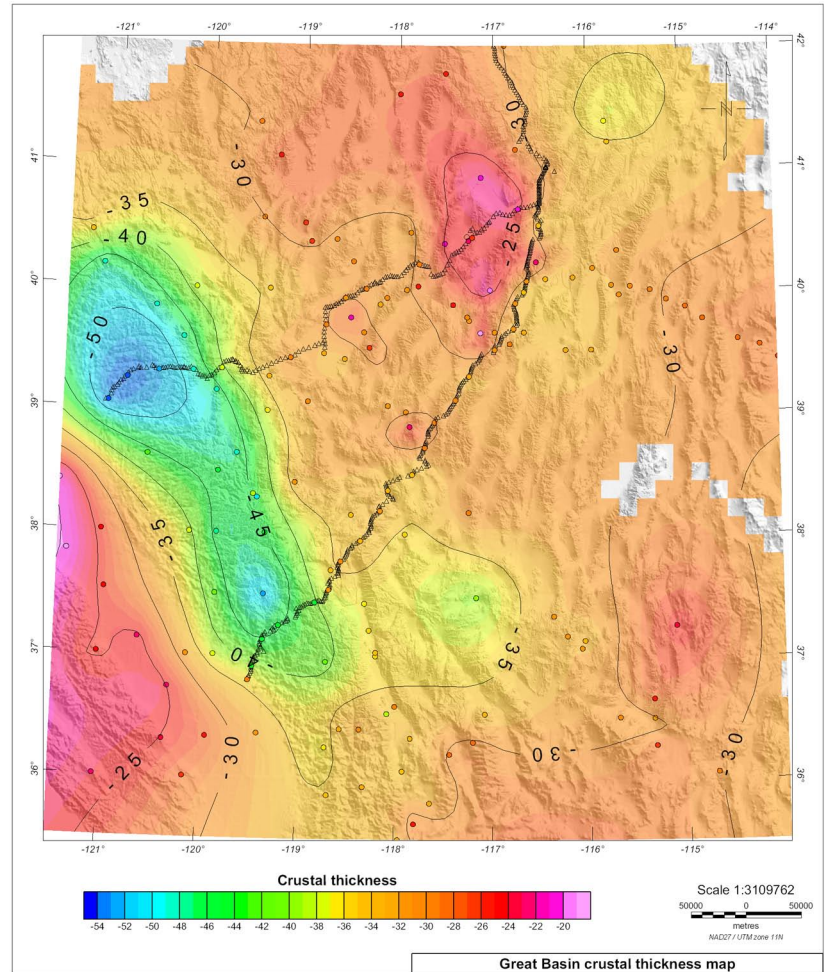


Figure 3. Crustal thickness map for the Great Basin and surrounding region. Colors are contoured at every 1 km change in thickness and the black lines are contoured at every 5 km. Colored circles indicate the location of data and the small black triangles denote the location of the NWL and INC experiments (also see Figure 1).

assume that the Pn, uppermost-mantle, velocity changes at the eastern boundary of the root where compositional and/or temperature changes are likely to occur. On the north side of Barrick GoldStrike, where we have no Pn first arrival picks, we have extrapolated the Moho to a depth of 31 km based on the recent refraction experiment of Lerch et al. (2007). The refraction experiment by Lerch et al. (2007) extends west of INC and across northern Washoe and Pershing Counties, Nevada, revealing approximately 31 km-thick crust that is not any thicker (as previously thought). On the western side of the Sierra Nevada, within the Great Valley, we have again extrapolated the Moho, this time to a depth of 28-30 km based on refraction experiments within the Great Valley and coast ranges (Fliedner et al., 2000; Mooney and Weaver, 1989).

3. Updated Crustal Thickness Map

We have compiled a contoured crustal thickness map (Figure 3) based on the NWL (Louie et al., 2004) and the updated INC seismic refraction results presented in this paper (Figure 2), as well as previous geophysical studies within the Great Basin. This paper presents an updated version of the crustal thickness map for the Great Basin. We extracted information concerning crustal thickness from the crust-mantle (Moho) refraction velocities and arrival times, cross-over distances, and from models presented in the literature. A kriging algorithm was used to interpolate a crustal thickness map. The colors on the map are contoured every 1 km, and the black outlines are for every 5 km, showing a more general trend. In addition, a preliminary version of this map was used in assessing areas for future refraction experiments, such as our NNUT experiment completed in August 2005 (Figure 1). On a first-order basis, the contoured crustal thickness data correlate well with heat flow observations for the Great Basin (Coolbaugh et al., 2005; Heimgartner et al., 2006).

The updated or preferred model eliminates most of the discrepancies between adjacent points, but leaves results in places where no other data exist. Combining these data sets allows better visualization of key features, permits discrepancies to be better addressed, and evaluates how well the proposed models fit the actual data. There are several key features to note: one, the crustal root is centered west of the Sierra's topographic crest; two, the root deepens north of Fresno, California; three, there is an abrupt step of approximately 18 km at the Sierra-Basin and Range transition; four, normal, 30-km thick Basin and Range crust occurs adjacent to the root on its east side; and five, the thin area of crust near Battle Mountain, Nevada may be larger than 100 km in area based on the compilation of the NWL and previous work in the area.

Conclusions

The Northern Walker Lane, Idaho-Nevada-California and Northern-Nevada-Utah transects are contributing to our efforts to understand regional geophysical properties and their relationships to the occurrence of geothermal resources. Our refraction experiments will help provide a more accurate crustal model for the northern Great Basin where little previous seismic refraction control exists. A detailed resolution of crustal thickness makes it possible to quantitatively evaluate the role that crustal-scale processes play in controlling the location of geothermal systems, and that evaluation is in progress.

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